



# Geological and geophysical characterization of the southeastern side of the High Agri Valley (southern Apennines, Italy)

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**Abstract.** In the frame of a national project funded by Eni S.p.A. and developed by three institutes of the National Research Council (the Institute of Methodologies for Environmental Analysis, the Institute of Research for Hydrogeological Protection and the Institute for Electromagnetic Sensing of the Environment), a multidisciplinary approach based on the integration of satellite, aero-photogrammetric and in situ geophysical techniques was applied to investigate an area located in the Montemurro territory in the southeastern sector of the High Agri Valley (Basilicata Region, southern Italy).

This paper reports the results obtained by the joint analysis of in situ geophysical surveys, aerial photos interpretation, morphotectonic investigation, geological field survey and borehole data. The joint analysis of different data allowed us (1) to show the shallow geological and structural setting, (2) to detect the geometry of the different lithological units and their mechanical and dynamical properties, (3) to image a previously unmapped fault beneath suspected scarps/warps and (4) to characterize the geometry of an active landslide affecting the study area.

gered (Guzzetti et al., 1996; Martino et al., 2004). So, the definition of the litho-structural setting is one of the basic aspects in landslide hazard assessment.

The study of landslide areas requires a multidisciplinary approach based on the integration of different techniques, providing a significant number of information, that can help to overcome the intrinsic limitations and drawbacks of each technique. The best choice is represented by the combined application of satellite, airborne and ground-based techniques (Perrone et al., 2006; de Bari et al., 2011). The first two techniques allow obtaining a synoptic view of the investigated area providing information from a small to a medium spatial scale. They can provide indication on superficial characteristics of the investigated area like the presence of geological structures and geomorphological features, land use, vegetation cover, etc. (Simoniello et al., 2008; Roering et al., 2009; Strozzi et al., 2010). Ground-based techniques give direct and indirect information on the subsurface characteristics in a specific point of the area (Petley et al., 2005; Marcato et al., 2012).

In situ non-invasive and low-cost geophysical techniques are widely employed in the subsurface investigations (Improta et al., 2010; Siniscalchi et al., 2010; Tropeano et al., 2013; Perrone et al., 2014). In particular, the electrical resistivity tomography (ERT) and the horizontal-to-vertical spectral ratio (HVSr) techniques have been widely applied in seismic microzonation (Boncio et al., 2011; Moscatelli et al., 2012), seismotectonic and geomorphological studies (Giocoli et al., 2008), to characterize the seismogenic fault systems (Galli et al., 2014; Giocoli et al., 2011), to reconstruct

## 1 Introduction

The origin and development of landslides may be influenced by several geologic, biologic, climatic and topographic factors such as rock strength, precipitation rate and intensity, hillslope gradient, vegetation type and density, bedrock fracture density, uplift rate, etc. The tectonic lineaments may represent preferential plans along which landslides can be trig-

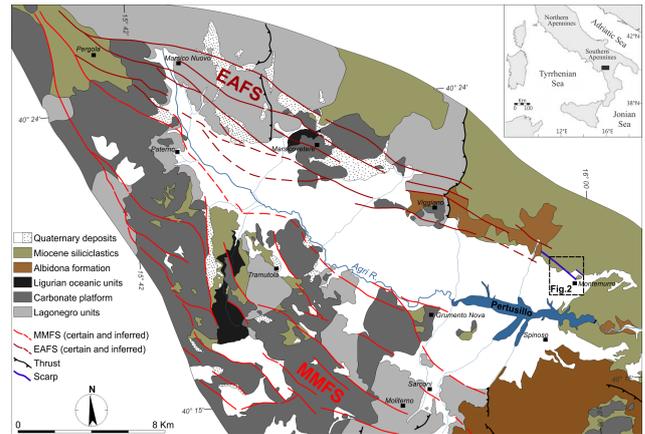
the geometry and the mechanical properties of the superficial litho-stratigraphic units (Mucciarelli et al., 2011; Albarello et al., 2011; Gallipoli et al., 2013), to delineate landslide body (Gallipoli et al., 2000; Perrone et al., 2004; Mainsant et al., 2012) and to study coseismic liquefaction phenomena (Giocoli et al., 2014).

In the frame of the project “Development and integration of innovative techniques of Earth Observation for landslides monitoring in a test site of the Agri Valley”, funded by Eni S.p.A., a multidisciplinary approach based on the joint application of satellite and ground-based Differential Interferometric Synthetic Aperture Radar (DInSAR) technique, aero-photogrammetric analysis and in situ geophysical measurements was employed to study an area located in the Montemurro territory in the southeastern sector of the High Agri Valley (Basilicata Region, southern Italy). This area represents a high seismogenic and hydrogeological hazard area, as testified by the *M* 7, 1857 Basilicata earthquake that produced a big landslide and 5000 casualties (Almagià, 1910).

This paper reports the results obtained by the joint analysis of in situ geophysical surveys, aerial photo interpretation, morphotectonic investigation, geological field survey and borehole data. The joint analysis of different data allowed us to (1) show the shallow geological and structural setting, (2) detect the geometry of the different lithological units and their mechanical and dynamical properties, (3) image a previously unmapped fault beneath suspected scarps/warps and (4) characterize the geometry of an active landslide that is still causing damage to structures and infrastructure.

## 2 Geological setting

The High Agri Valley (HAV) is a Quaternary NW–SE trending intermontane basin located in the axial zone of the Southern Apennines thrust belt (southern Italy) (Fig. 1). Brittle tectonics has strongly controlled the formation and evolution of the HAV up to the present. The deformation is testified by seismic activity and by loose slope deposits and paleosols involved in faulting in the last 40 ka (Giano et al., 2000; D’Addezio et al., 2006). The HAV was hit by the *M* 7, 1857 Basilicata earthquake. A seismogenic structure capable of producing large events is alternatively associated with: (1) the Monti della Maddalena Fault System (MMFS) (Pantosti and Valensise, 1988; Maschio et al., 2005; Improta et al., 2010) and (2) the Eastern Agri Fault System (EAFS) (Benedetti et al., 1998; Cello et al., 2003; Giano et al., 2000). The MMFS bounds the HAV to the SW and runs for about 25 km between Pergola and Moliterno villages, whereas the EAFS, which is associated with mature fault-line scarps, bounds the HAV to the NE. No detailed definition exists on the length of the EAFS that, based on mapped faults between Pergola and Viggiano villages, is about 25 km (Cello et al., 2003; Maschio et al., 2005).



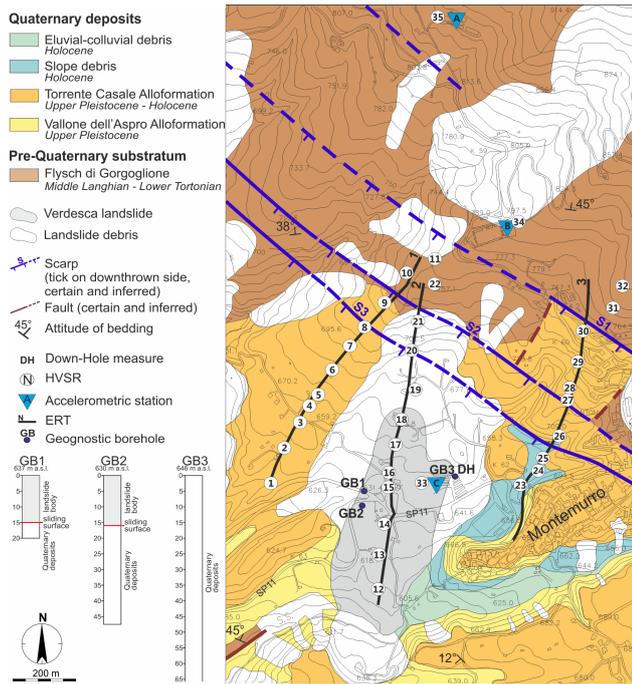
**Figure 1.** Schematic geological map of the High Agri Valley.

The study area falls in the southeastern part of the HAV, in the territory of Montemurro (Fig. 2). The pre-Quaternary substratum of this area consists of Tertiary siliciclastic sediments of the Gorgoglione Flysch (GF) that crop out mainly in the northern sector of Montemurro. Downslope, the Gorgoglione Flysch is covered by Quaternary continental deposits (QD) related to both Vallone dell’Aspro and Torrente Casale Alloformations (Zembo et al., 2010).

The morphology of the Montemurro territory is the result of different concomitant factors, such as the Quaternary regional uplift and the consequent fluvial erosion. Land modelling is also controlled by lithology and the structural setting. General features of the area are the occurrence of steep slopes, encased streams within narrow and deep land incisions and abrupt acclivity change, which could correspond to tectonic structures or lithological variations.

The Montemurro territory is affected by extensive hydrogeological instability phenomena. Active and quiescent landslides involve both the QD and the GF. In particular, our analysis shows an active landslide, which is still causing damage to residences, anthropic buildings and, especially, to the main road (SP11) leading to the urban area of Montemurro. The identified active landslide was termed Verdesca landslide. It is approximately 750 m long, 320 m wide and extends between 675 and 590 m a.s.l. and has a mean inclination of about 7° (Fig. 2).

Our surficial investigations (e.g. field geological survey, aerial photo interpretation) show also rectilinear NW–SE trending and SW-facing scarps (S1–S3 in Fig. 2). These scarps are perpendicular to the strike of the streams and they are aligned to the main trend of the EAFS, which was clearly identified between the villages of Pergola to the north and Viggiano to the south (Cello et al., 2003; Maschio et al., 2005). Thus, we suspect that the NW–SE scarps observed in the field could represent the surface expression of an additional strand of the EAFS. In order to unravel this uncertainty, in situ geophysical surveys were carried out.



**Figure 2.** Simplified geological map of the study area showing the distribution of the NW–SE scarps observed in the field, the Verdesca landslide, the location of the surveys and the three simplified borehole logs.

### 3 Geophysical surveys

Geophysical methods provide an efficient tool for imaging the subsurface. The suitability of a particular geophysical technique or a combination of them mainly depends on the physical property contrast between the target and the host material, depth extent of the target, and the nature and thickness of the overburden. Generally speaking, the investigations of as many properties as possible by various geophysical methods enable a double-check of results and enhance the reliability of interpretation. In this paper, we focused our attention on two complementary techniques, the electrical resistivity tomography (ERT) and the horizontal-to-vertical spectral ratio (HVSr) by earthquakes and ambient noise measurements. These two techniques are among the most non-invasive, fastest and cheapest geophysical methods. All the surveys were carried out in an area located in the Montemurro territory between April 2010 and May 2012 (Fig. 2) and were aimed at: (1) imaging the structural and sedimentary setting, (2) verifying the nature of the NW–SE scarps in order to interpret them either as a tectonic landform (i.e. fault scarp) or as a geomorphic feature (i.e. erosional scarp) and (3) characterizing the Verdesca landslide (e.g. geometry, thickness).

The ERT and HVSr results were supported by interpretation of aerial photos, morphotectonic investigation, geological field survey and were compared with the stratigraphical

data obtained by three geognostic wells reported in Fig. 2 as GB1 (20 m deep), GB2 (48 m deep) and GB3 (66 m deep) (Fig. 2).

#### 3.1 Electrical resistivity tomography

Three NE–SW ERT were carried out across the NW–SE scarps observed in the Montemurro territory (S1–S3 in Fig. 2). Only the ERT2 profile runs across the Verdesca landslide.

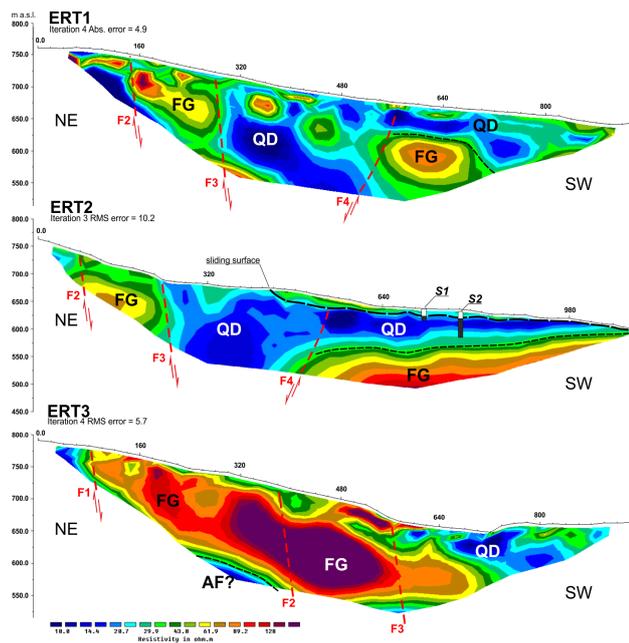
To acquire resistivity data, a resistivimeter Syscal R2 (Iris Instruments) connected to a multielectrode system with 48 electrodes, arranged according to the Wenner–Schlumberger array and spaced 20 m, was used. The ERT1 and ERT3 were performed along 940 m long profiles and have reached an investigation depth of about 150 m. ERT2 was realized along a 1100 m long profile using roll-along technique and reaching the same investigation depth.

Apparent resistivity data were inverted using the RES2DINV software (Loke, 2001). Each resistivity model was chosen at the interaction after which the root mean square (rms) did not change significantly. This occurred between the third and fourth interactions. In all cases, the rms error was less than 11 % and the resistivity values range from 10 to more than 128  $\Omega\text{m}$  (Fig. 3).

Taking into account the data gathered through geological surveys, aerial photo interpretation and exploratory boreholes, we were able to calibrate the resistivity models and to directly correlate resistivity values with lithostratigraphic characteristics. Thus, the higher resistivity values ( $> 35 \Omega\text{m}$ ) can be associated with the Gorgoglione Flysch (GF) and the lower resistivity values ( $< 30 \Omega\text{m}$ ) are related to the Quaternary continental deposits (QD) or to the Albidona Formation (AF).

From a structural viewpoint, the major features of the resistivity models are the sharp lateral variations of resistivity F1, F2 and F3 (Fig. 3), which fit at the surface with the NW–SE scarps S1, S2 and S3 (Fig. 2), respectively. In addition, these scarps are aligned to the main trend of the EAFS and are perpendicular to the strike of the streams. Thus, we can speculate that the scarps are a surface expression of an additional strand of the EAFS. The identified strand was termed Montemurro Fault (MF). The resistivity models (ERT1 and ERT2) also show a NE dipping splay (F4) of the MF with no geomorphic signature (e.g. scarp). F4 was inferred also by Stabile et al. (2014) based on the joint analysis of seismicity data, geological observations, fluid injection data, stratigraphic log of the Costa Molina 2 injection well, and ERT survey.

Overall, our investigations show that the MF consists of at least four sub parallel splays (F1, F2, F3 and F4). In particular, between F3 and F4, the ERT1 and ERT2 show a tectonic depression. Here the ERT do not allow us to investigate the substratum (GF) that is more than 150 m depth. In the upper part of the ERT2, between 430 and 1100 m, the moderate re-



**Figure 3.** Resistivity models (see Fig. 2 for the location of ERT profiles).

sistivity values (25–70  $\Omega\text{m}$ ) are associated with the body of the Verdesca landslide. The estimated thickness of the slide deposits, as derived from borehole data (GB1–GB3 in Fig. 2) and resistivity contrast observed in the ERT2, varies from less than 10 up to 30 m.

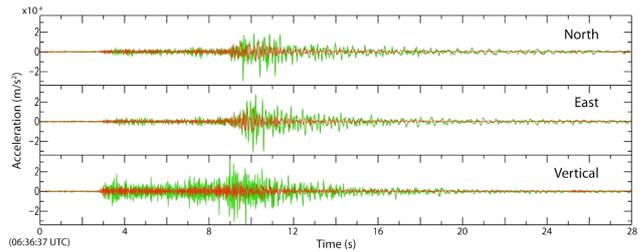
### 3.2 Earthquake and ambient noise horizontal-to-vertical spectra ratio

In the study area we carried out 35 ambient noise measurements, each one with duration of 12–16 min, with a digital tri-directional tromometer, a high-resolution seismometer whose 24-bit dynamic is aimed at the very low amplitude range, starting from the Brownian motion of the sensor (Tromino<sup>®</sup> Micromed).

Moreover, in February 2011 we installed two ETNA accelerometers: one in the landslide body (accelerometric station C, Fig. 2) and the other one outside the landslide, on the sandstones of the Gorgoglione Flysch (GF, bedrock) (accelerometric station A, Fig. 2). Due to logistic problems, on October 2011 the accelerometer installed on the bedrock was moved from the site A to the site B (see Fig. 2). During all the period we recorded 44 local and regional earthquakes with local magnitude ( $M_l$ ) ranging between 1.1 and 6.0 (regional event).

Figure 2 shows the location of the seismic ambient noise measurement points, the accelerometric stations and the available down-hole survey.

Table 1 reports the list of earthquakes recorded by accelerometers with their principal seismological parameters.



**Figure 4.** Comparison between waveforms of the  $M_l = 2.7$  18 February 2011 Potenza earthquake recorded inside the landslide body (green curve, accelerometric station C) and on the bedrock (red curve, accelerometric station A).

In particular, the last column reports indication of stations that recorded the earthquakes acquired during February 2011 to March 2012. Figure 4 shows the waveforms of the  $M_l = 2.7$  Potenza earthquake of 18 February 2011 recorded inside the landslide body (green curve) and on the bedrock (red curve). It is possible to observe the clear site amplification inside the landslide body with respect to the site A (bedrock), with evident higher peak ground acceleration amplitudes and longer signal duration.

The earthquakes and ambient noise recordings have been analysed by the HVSR techniques (Lermo and Chavez-Garcia, 1993; Castro et al., 1996). The HVSR functions were calculated by averaging the HVSR obtained by dividing the signal into non-overlapping windows of 20 s. Each window was detrended, tapered, padded, FF-Transformed and smoothed with triangular windows with a width equal 5 % of the central frequency. The average was used to combine E–W and N–S components in the single horizontal (H) spectrum. Average single-component spectra were obtained from the same procedure. For each HVSR curve the relative  $\pm 2\sigma$  confidence interval is given. Some authors suggest that transient can affect estimates of fundamental frequency of soils, but a simple variation of amplitude never caused this problem according to Parolai and Galiana-Merino (2006) and Mucciarelli (2007). The ambient noise recordings and HVSR analysis have been estimated according SESAME (Chatelain et al., 2008) and Albarello et al. (2011) guidelines.

In order to validate the ambient noise HVSRs, we compared ambient noise HVSR related to the sites 33, 34 and 35 with those obtained from earthquake recordings at the sites A, B and C (Fig. 2). In particular, the ambient noise measurement related to site 35 was carried out close to station A, the one related to site 34 near to station B, and the ambient noise recording of site 33 was performed in correspondence with station C. Figure 5a and b shows flat HVSRs for both components of the accelerometric station A, which was installed on the bedrock; on the other hand, the HVSR of the station B has an evident resonance peak between 2 and 3 Hz (Fig. 5c and d). Finally, HVSR of the station C shows two clear resonance peaks at 1.5 and 1.8 Hz, respectively (Fig. 5e and f).

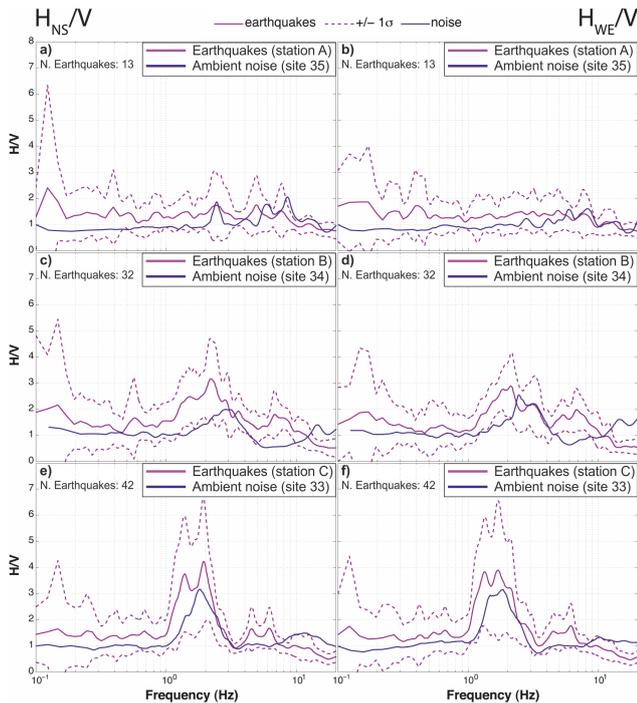
**Table 1.** List of earthquakes recorded from February 2011 to March 2012 by the accelerometric stations A, B, and C installed in Montemurro.

Data	Origin time (yyyy-mm-dd)	Lat (° N)	Long (° E)	Depth (km)	Place	Mag.	Stations
2011-02-18	06.36.32	40.626	15.804	11.1	Potenza	2.7 $M_1$	A and C
2011-03-10	–	–	–	–	Local event	–	A and C
2011-03-11	–	–	–	–	Local event	–	A and C
2011-04-01	13.29.11	35.540	26.630	60.0	Crete (Greece)	6.0 $M_w$	A and C
2011-04-02	16.25.38	39.921	16.006	6.8	Pollino	2.7 $M_1$	A and C
2011-04-12	03.32.55	41.663	16.117	10.1	Gargano	3.9 $M_1$	A and C
2011-04-25	15.32.07	40.310	16.268	26.0	Lucanian Apennine	1.9 $M_1$	A and C
2011-07-07	–	–	–	–	Local event	–	A and C
2011-07-12	–	–	–	–	Local event	–	A and C
2011-08-05	15.43.54	40.479	15.612	8.6	Le Murge	3.1 $M_1$	A and C
2011-09-20	03.08.02	40.243	15.674	–	Diano Valley	2.4 $M_1$	A and C
2011-09-21	03.26.21	40.247	15.671	9.3	Diano Valley	2.3 $M_1$	A and C
2011-11-23	14.12.33	39.915	16.018	6.3	Pollino	3.6 $M_1$	B and C
2011-11-24	02.13.05	39.904	16.014	8.5	Pollino	2.8 $M_1$	B and C
2011-11-30	06.21.23	40.339	15.992	10.0	Viggiano	1.0 $M_1$	B and C
2011-11-30	–	–	–	–	Local event	–	B and C
2011-11-30	21.07.21	39.921	16.021	8.4	Pollino	2.5 $M_1$	B and C
2011-12-01	14.01.20	39.932	16.002	6.8	Pollino	3.3 $M_1$	B and C
2011-12-01	18.02.47	39.921	16.015	9.1	Pollino	2.5 $M_1$	B and C
2011-12-01	–	–	–	–	Local event	–	B and C
2011-12-02	09.31.10	39.917	16.011	8.9	Pollino	2.6 $M_1$	B and C
2011-12-02	19.38.44	39.920	16.005	9.2	Pollino	2.5 $M_1$	B and C
2011-12-02	21.25.38	39.914	16.013	8.8	Pollino	3.2 $M_1$	B and C
2011-12-03	06.16.24	39.939	16.015	8.0	Pollino	2.5 $M_1$	B and C
2011-12-04	–	–	–	–	Local event	–	B and C
2011-12-08	04.58.34	39.937	16.017	7.7	Pollino	2.5 $M_1$	B and C
2011-12-10	14.53.25	39.915	15.983	8.6	Pollino	2.6 $M_1$	B and C
2011-12-15	–	–	–	–	Local event	–	B and C
2011-12-17	23.20.15	39.373	16.175	5.5	Valle del Crati	3.4 $M_1$	B and C
2011-12-19	–	–	–	–	Local event	–	B and C
2011-12-20	10.38.07	40.474	15.792	10.2	Lucanian Apennine	1.9 $M_1$	B and C
2011-12-23	21.54.09	39.942	16.021	9.6	Pollino	2.3 $M_1$	B and C
2011-12-24	20.17.50	39.914	16.037	8.3	Pollino	3.3 $M_1$	B and C
2012-01-12	20.28.53	39.932	16.014	8.6	Pollino	2.5 $M_1$	B and C
2012-01-15	06.32.27	40.246	15.902	9.2	Agri Valley	2.4 $M_1$	B and C
2012-01-15	12.35.36	40.243	15.912	9.5	Agri Valley	1.7 $M_1$	B and C
2012-01-15	22.36.18	39.912	16.003	7.4	Pollino	2.6 $M_1$	B and C
2012-01-16	01.42.41	40.247	15.910	9.4	Agri Valley	1.5 $M_1$	B and C
2012-01-22	10.04.02	39.288	15.599	22.1	West Calabrian Coast	3.5 $M_1$	B and C
2012-01-24	07.42.56	40.321	15.983	9.8	Agri Valley	1.1 $M_1$	B and C
2012-01-25	–	–	–	–	Local event	–	B and C
2012-03-09	23.26.46	40.168	15.992	10.1	Agri Valley	1.9 $M_1$	B and C
2012-03-12	02.10.09	40.469	16.076	28.4	Lucanian Apennine	2.7 $M_1$	B and C
2012-03-17	04.54.51	40.537	15.892	23.0	Lucanian Apennine	2.2 $M_1$	B and C

In all the three cases considered, the ambient noise HVSRs (blue lines in Fig. 5) are in agreement with the earthquake HVSRs (pink lines in Fig. 5). The availability of accelerometric recordings on a bedrock site (station A) allowed us to also perform standard spectral ratio (SSR) analyses between waveforms recorded on the landslide body and those recorded at the same time on the reference site. The SSR

analysis confirms the resonance peak at about 1.8 Hz in the landslide body.

In order to attribute stratigraphic and mechanic meanings to resonance peak at 1.8 Hz revealed in the earthquake/ambient noise HVSRs of the station C and site 33 (landslide body), numerical analysis was performed using the model proposed by Castellaro and Mulargia (2009), the 30 m depth down-hole test (Fig. 6a) and the mechanical and



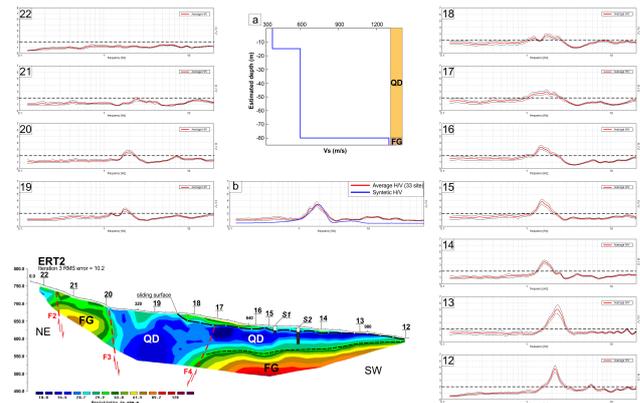
**Figure 5.** HVSR curves obtained from earthquake data (pink curves) and the ambient noise measurement (blue curves) for both horizontal components ( $H_{NS}/V$ , left panels;  $H_{WE}/V$ , right panels).

seismic data available in the area. The stratigraphic sequence is characterized by the Quaternary continental deposits (QD) (see also GB3 in Fig. 2) with  $S$  wave velocity ( $V_s$ ) increasing with depth from  $380 \text{ m s}^{-1}$  in the first 15 m to  $600 \text{ m s}^{-1}$  at depth  $> 15 \text{ m}$ . At the bottom, the  $V_s$  of the Gorgoglione Flysch (GF) is more than  $800 \text{ m s}^{-1}$  (seismic bedrock) (Fig. 6a).

The inversion model provided a synthetic HVSR in good agreement with the experimental one (Fig. 6b). The resonant peak at about 1.8 Hz is related to the depth of the seismic impedance contrast between the GF and the overlying QD with 80 m thick (resonant layer).

Subsequently, we have extended the seismic ambient noise measurements along all the ERT profiles in order to improve the interpretation of the results by using the joint analysis. Figure 6 shows ambient noise average HVSRs superimposed along ERT2. The HVSRs 12 and 13 show a clear resonance peak at 2.2 Hz. This peak shifts to lower frequency (1.5 Hz) for the site 14 to the site 15 (central part of the profile); the northern HVSRs (sites 21 and 22) show flat curves because they lay on the outcropping seismic bedrock (Gorgoglione Flysch). Thus, the results summarized in Fig. 6 point out that the fundamental frequency is related to the depth variation of the seismic impedance contrast between the high-resistive GF and the overlying low-resistive QD.

Finally, it is possible to observe in Fig. 6 a little lowering of the average HVSR in correspondence of 1.7 Hz at sites 18–21. The disturbance is attributable to the presence of



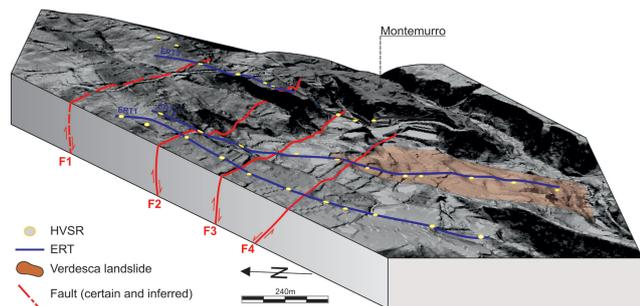
**Figure 6.** Comparison between the ambient noise HVSRs and the ERT2. 12 to 22 indicate the HVSR curves from seismic noise measurements. (a) Value of  $V_s$  used to constrain the H/V curve. (b) Comparison between the synthetic HVSR curve and the ambient noise HVSR curve at the site 33.

in-service wind turbine towers (Saccorotti et al., 2011) near those sites, and it is principally evident on the HVSR obtained for the E–W component. Since such frequency is near the resonance peak, in these four sites we reconstructed the HVSR function by considering only the N–E component.

#### 4 Discussions and conclusions

In the frame of the project “Development and integration of innovative techniques of Earth Observation for landslides monitoring in a test site of the Agri Valley” funded by Eni S.p.A. a multidisciplinary approach based on the integration of satellite, aero-photogrammetric and in situ geophysical techniques was applied to investigate an area located in the Montemurro territory in the southeastern sector of the HAV. In this paper we focused our attention on the results obtained by the joint application of two in situ geophysical techniques, the ERT and the earthquake and ambient noise HVSR. The interpretation of results was supported by morphotectonic investigation, interpretation of aerial photos, geological field survey and borehole data. In particular, the joint application of these two in situ geophysical techniques allowed us to infer a good correspondence between the resistivity contrast and the seismic impedance interface. Both interfaces coincide with the boundary between the GF and the overlying QD characterized by different resistivity and mechanical behaviour.

Overall, the joint analysis of aerial photo interpretation, morphotectonic investigation, geological field survey, borehole data, ERT and HVSR allowed us to: (1) image the shallow geological and structural setting, (2) detect the geometry of the different lithological units and their mechanical and dynamical properties, (3) image the geometry and architecture of the previously unmapped Montemurro Fault (MF)



**Figure 7.** 3-D surface digital elevation model of the Montemurro area showing the faults and the location of ERT profiles and HVSR.

(Fig. 7) and (4) verify the nature of the NW–SE scarps and to interpret them as the surface expression of the MF.

The HVSR technique does not show resonance peaks related to the depth of the sliding surface of the Verdesca landslide due to a negligible impedance contrast between the QD and the material of the Verdesca landslide. Thus, only ERT and exploratory boreholes were integrated to: (1) characterize the geometry of the Verdesca landslide, (2) locate the sliding surface and (3) estimate the thickness of the landslide body.

We think that the results presented in this study provide new basic data for hydrogeological and seismic hazard assessment of the HAV. In particular, the HAV is one of the areas of Italy with the highest seismogenic potential. In addition, the seismic risk is amplified by the presence of two important human activities, namely the exploitation of the largest productive on-shore oil field in West Europe, managed by the Eni company, and the artificial Pertusillo water reservoir (Stabile et al., 2014a, b). The recognition of the Montemurro Fault, which increases longitudinal extent of the EAFS previously mapped by Cello et al. (2003) and Maschio et al. (2005), may indeed have significant implications for the local seismic hazard. However, our data do not show whether the EAFS is the seismogenic structure of the High Agri Valley, which remains undefined.

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