Decrease in the electric intensity of VLF/LF radio signals and possible connections

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Abstract. In February 2002, a receiver was put into operation at the Department of Physics of Bari University (Southern Italy) to record VLF-LF radio signals. The intensity and the phase of the signals transmitted by GB (f=16 kHz, United Kingdom), FR (f=20.9 kHz, France), GE (f=23.4 kHz, Germany), IC (f=37.5 kHz, Island) and IT (f=54 kHz, Sicily, Italy) has been monitored with a 5 s sampling rate. The intensity raw data averaged over 10 min, from February 2002 to April 2006, have been analysed. Several decreases of the electric field intensity of the radio signals with a duration of some days were revealed, generally occurring in not concomitant periods. The GE signal decreases systematically in winter and summer each year and so, it could be supposed that such decreases are related to the transmitter. On the contrary, all the other decreases pointed out are sporadic and the previous justification does not seem realistic. On the time occurrence of these decreases, the geomagnetic activity, the meteorological conditions in the receiver area and the regional seismic activity were investigated. The main result is that, generally, a pre or post seismic effect seems to give the most convincing justification. The effect seems to appear for earthquakes with magnitude $M \geq 4.3$, when the epicentres are within the third Fresnel zone of the radio signals or near enough to some radio path.

1 Introduction

For many years research into the interaction between seismic activity and disturbances in radiobroadcasts has been carried out. One of the first results was obtained using 18 MHz receivers on the occasion of the great ($M=8.5$) Chilean earthquake of 22 May 1960 (Warwick et al., 1982). The receivers were part of a network used for studying cosmic noise. Later, pre-seismic disturbances in VLF (3–30 kHz) – LF (low band, 30–60 kHz) radio signals, revealed by ground based measurements, have been presented by Japanese and Russian researchers (Hayakawa and Sato, 1994; Hayakawa et al., 1996; Hayakawa et al., 2002; Hayakawa et al., 2006; Molchanov and Hayakawa, 1998; Rozhnoi et al., 2005; Rozhnoi et al., 2006). Moreover, pre-seismic disturbances on LF (high band, 150–250 kHz) radio broadcasts pointed out by ground receivers were proposed by Italian researchers (Biagi et al., 2001a, b; Biagi and Hayakawa, 2002; Biagi et al., 2005; Biagi et al., 2006). VLF-LF (low band) radio signals radiated by powerful transmitters located on the ground can be received not only by ground receivers but also on board of satellites. Recently, some possible seismic disturbances revealed by VLF radio signals measured on board of the French DEMETER satellite, were presented by Molchanov et al. (2006).

2 Equipment and data collecting

On February 2002, in the framework of a scientific cooperation among Japanese, Russian and Italian teams, a receiver able to measure the electric field strength (intensity)
and the phase of five VLF-LF (low band) radio signals was put into operation at Bari (South Italy). The receiver is the Japanese OmniPAL model connected to a computer by a digital card and it is located in the Department of Physics of Bari University. The electric antenna and the GPS sensor are installed on the roof of the building. On the basis of the best reception at the receiver and taking into account the most convenient radio paths in relation to the seismic activity, the following transmitters were selected: GB \((f=16\text{ kHz}, \text{ United Kingdom})\), FR \((f=20.9\text{ kHz}, \text{ France})\), GE \((f=23.4\text{ kHz}, \text{ Germany})\), IC \((f=37.5\text{ kHz}, \text{ Iceland})\) and IT \((f=54\text{ kHz}, \text{ Sicily, Italy})\). The different radio-paths are indicated.

![Map showing the location of the receiver (Bari) and of the VLF/LF transmitters: GB \((f=16\text{ kHz}, \text{ United Kingdom})\), FR \((f=20.9\text{ kHz}, \text{ France})\), GE \((f=23.4\text{ kHz}, \text{ Germany})\), IC \((f=37.5\text{ kHz}, \text{ Iceland})\) and IT \((f=54\text{ kHz}, \text{ Sicily, Italy})\). The different radio-paths are indicated.](image)

3 Data analysis

In order to manage the huge amount of data (at 5 s sampling rate) we analysed the time series of 10 min means (one datum each 10 min averaged between the raw data 5 min before and 5 min after the datum). Figure 2a shows the raw data (bottom) and the averaged data (top) related to the GE transmitter during October 2005. The radio signals show standard features with high intensity value at night time and low one at day time. It can be observed that the averaging procedure does not affect the main features of the signal. The relative trends of the five radio signals were analysed to search for irregularities. The main result of this analysis was the finding of several decreases in the intensity of the radio signals with a duration of some days. This phenomenology stood up clearly by examining the trends smoothed by a running adjacent averaging over 5 days. As an example the Fig. 2b, at the bottom, shows the smoothed trend of the GE radio signal. Many clear decreases of the signal intensity appear, but it is evident that in many cases the effect is a periodic one, i.e. in December–January and in June–July. It is reasonable to assume that these decreases are not connected with natural phenomena but with systematic reductions of the radiated power by the transmitter. So, these decreases were not considered in our analysis and the trend at the top of the Fig. 2b represents the effective trend of the GE signal that we have examined. All the other decreases pointed out in the various radio signals are sporadic and have been considered. Generally, their occurrence is in not concomitant periods.

In the past, some influence of the geomagnetic activity and of the meteorological conditions on the propagation of the VLF-LF radio signals has been presented (Kikuchi, 1981; Kleimenova et al., 2004; Rozhnoi et al., 2006b). Then, possible effects of the seismic activity occurred in the area of the radio paths were pointed out (Hayakawa et al., 1996; Molchanov and Hayakawa, 1998; Rozhnoi et al., 2005; Rozhnoi et al., 2006a).

So, in order to give explanation of the previous decreases of the radio signals intensity, for each case the geomagnetic activity, the meteorological conditions in the receiver area and the seismic activity was investigated. For the geomagnetic activity the indexes \(K_p\) and \(D_{st}\) were considered; for the meteorological conditions the data of stations located in Bari were used and the air temperature, pressure and wind velocity were considered as the most indicative parameters (Rozhnoi et al., 2006b); for the seismic activity the earthquakes with \(M>4.0\) which occurred within the third Fresnel zone of each radio signal were considered. The quoted zone is considered to be the most sensitive to the seismo-atmospheric effects (Hayakawa et al., 1996; Molchanov and Hayakawa, 1998; Rozhnoi et al., 2005).

4 Results

Figure 3a shows the data measured during April 2002, when a substantial decrease was observed in the IT radio signal. On 17 April 2002 three earthquakes with \(M=3.0–4.3\) occurred in the gulf of Taranto and an aftershock (\(M=3.4\)) occurred on 27 April. The epicentres are shown in the Fig. 3b.
Fig. 2. (a) At the bottom, raw data on time series of the intensity of GE radio signal collected during October 2005. The trend at the top is the averaged 10 minutes time series. (b) The trend at the bottom is the averaged 10 min time series of the raw intensity of GE radio signal, smoothed by a running adjacent averaging over 5 days. At the top the same trend, where the systematic decreases in December–January and June–July are removed, is shown.

Fig. 3. (a) During April 2002, from the top reading downwards: $k_p$ geomagnetic index, $D_{st}$ geomagnetic index, air temperature, wind velocity, pressure and the averaged 10 min intensity of IT radio signal. The red line indicates the time occurrence of the main ($M=4.3$) earthquake of the gulf of Taranto. (b) Map showing the location of the earthquakes occurred in the gulf of Taranto during April 2002. The IT radio path is indicated.
Fig. 4. (a) During December 2003–January 2004, from the top reading downwards: $k_p$ geomagnetic index, $D_{st}$ geomagnetic index, air temperature, wind velocity, pressure and the averaged 10 minutes intensity of the GB radio signal. The red line indicates the time occurrence of the main ($M=4.3$) earthquake of the southern Apennines sequence. (b) Map showing the location of the earthquakes occurred in the southern Apennines at the end of December 2003. The GB radio path is indicated.

Fig. 5. (a) In the period 6 February–20 March 2004, from the top reading downwards: $k_p$ geomagnetic index, $D_{st}$ geomagnetic index, air temperature, wind velocity, pressure and the averaged 10 minutes intensity of FR and IT radio signals. The red line indicates the time occurrence of the earthquake ($M=4.6$) happened near the west coast of the Calabria region. (b) Map showing the location of the quoted earthquake occurred on 3 March 2004. The FR and IT radio paths are indicated.
Figure 4a shows the data measured during December 2003–January 2004, when a decrease was pointed out in the GB radio signal. On 30 December 2003 nine earthquakes with $M=2.1$–4.3 occurred in short sequence in the southern Apennines and two foreshocks ($M=3.1$) occurred on 28 December. The epicentres are shown in the Fig. 4b.

Figure 5a shows the data measured during 16 February–20 March 2004. Decreases appear in both IT and FR radio signals. At 3 March an earthquake with $M=4.6$ occurred near the west coast of the Calabria region, as it is shown in the Fig. 5b. In the same figure the IT and FR radio paths are indicated. As it concerns the FR radio signal it must be noted that the epicentre is out the third Fresnel zone.

Figure 6a shows the data measured during 8–12 November 2004. The decrease appearing in the intensity of the GE radio signal is the alone still present after removing the periodic ones (Fig. 2b). We haven’t found any reason to remove also this decrease because related to the transmitter. During this period, the seismic activity was not very intense within the third Fresnel zone or, more generally, within the radio path zone. The Fig. 6b shows the data measured during 6–8 December 2004, when a decrease was revealed in the IT intensity. Also in this case no significant earthquake occurred.

The most recent decrease was observed in the IC radio signal, as it appears in the Fig. 7a where the data measured during August-September 2005 are shown. On August 22 an earthquake with $M=4.7$ occurred at the Latium coast and on 8 September, an earthquakes with $M=4.4$ happened near the north-west border of Italy; these epicentres and the radio path are shown in the Fig. 7b. It must be noted that the epicentre of the earthquake at the Latium coast is out the third Fresnel zone of the radio path, but it was considered due to the strict connection of its location with the IC radio path represented by a well known tectonic line (called Anzio-Ancona) in central Italy. This line is indicated in red colour in the Fig. 7b.

5 Discussion

On the basis of the figures showed in the previous sections, it is possible to identify connections between the radio decreases and different kind of phenomenology that could be responsible of these anomalies. The most indicative causes of perturbation in the radio signal seem to be the seismic and the geomagnetic activity, and some meteorological disturbances. Anyway, if a correlation is claimed between the radio anomaly and the geomagnetic activity, they must happen contemporaneous; on the contrary, a temporal shift is reasonable if a correlation between the radio anomaly and the seismic activity is considered, so that both preseismic and postseismic effects can be revealed.

That being stated, looking at the figures, it is possible to note that there is an only case of clear connection between a
radio decrease and a contemporary increase in the geomagnetic activity, as shown in Fig. 6a, where the GE signal is presented. Instead, no clear correlation appears if we take into consideration radio decreases and variations in meteorological parameters, due to the shift in time, always existing in all the cases presented, between the two events.

Now, we can discuss the connection between radio signal disturbances and seismic activity.

Figure 3a shows an intensity decrease in the IT radio signal, lasting about 4 days and its connection with the earthquakes happened in the gulf of Taranto, indicated in the Fig. 3b, seems convincing. So, a radio precursor with a duration of three-four days is revealed. In the same way, the intensity decrease in the GB radio signal presented in the Fig. 4a can be connected with the earthquakes (southern Apennines) shown in the Fig. 4b, so that it appears as a precursor with a duration of about seven days. In this case also a post seismic effect with a similar duration (Fig. 4a) can be claimed. The situation presented in the Fig. 5a shows two different decreases in the FR and in the IT radio signals. The connection of the IT decrease with the earthquake \(M=4.6\) happened at the west coast of the Calabria region seems realistic. In such a case a postseismic effect with a duration of 9–10 days can be claimed. On the other hand, the connection of the FR decrease with the same earthquake is more doubtful, due to the location of the epicentre with respect to the FR radio path, but it could be considered and in this case it becomes a possible precursor with a duration of 10 days. Finally, a decrease of the IC signal is presented in the Fig. 7a. The connection with the earthquake \(M=4.4\) that happened near the north-west border of Italy is convincing and probably also a connection with the earthquake \(M=4.7\) occurred at the Latium coast can be claimed. In fact, the well known tectonic line Anzio-Ancona connects the location of this epicentre with the radio path as we noted in the previous section. So, probably, two different anomalies exist on the IC radio signal and they represent two precursors with a duration of 8–12 days. It is possible to note that, in this case, also the geomagnetic activity increases during the occurrence of the radio anomaly but with a shorter duration, so that some minor effects connected with this phenomenology can be claimed.

In the Fig. 6b a decrease of the IT radio signal is presented and this is the only case where no possible relations with any parameter has been found out; also the seismic activity inside and outside the relative Fresnel zone was extremely moderate.

So, during the period of the measurements, eight decreases of the intensity of VLF-LF radio signals were revealed. In six cases the connection with the seismicity seems the most
convincing cause, with some doubt concerning the FR decrease in Fig. 5a; in one case the connection with the geomagnetic activity is evident. Only in one case no connection was revealed. The “seismic cases” reveal precursors for five times and for one time only a postseismic effect.

In order to support the “seismic cases”, it must be noted that the events reported in the Fig. 5 and in the Fig. 7 were the largest earthquakes occurred within a radius of 100 km, during the whole period of the radio signals monitoring: therefore, they represent rare events. Unfortunately, a phenomenon very similar to the seismic sequence reported in the Fig. 3 happened after the failure of the antenna, in April 2006, so, a possibility of comparison was impossible, as well for the other (in addition to that shown in the Fig. 4) large earthquake ($M=5.4$, 31 October 2002), within a radius of 100 km, that happened in a period when the data were not collected, due to maintenance of the location where the receiving antenna is located (Biagi et al., 2004).

As it concerns the theoretical justification of the decreases of the radio signals intensity in connection with the seismicity, a possible model based on the acoustic gravity waves and the ionosphere turbulence was proposed by Molchanov et al. (2006).

6 Conclusions

A good possibility exists that the earthquakes with $M \geq 4.3$ can produce some decrease in the intensity of the VLF-LF radio signals, when the path is near enough to the epicentres. The decreases can represent a preseismic effect or a postseismic one.

The systematic measurement of the intensities of different VLF-LF radio signals in several points could increase the definition of the effect.

So, the starting up of networks for monitoring VLF-LF radio signals is hoped for.

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